# AFRICAN DUST DEPOSITION AND OCEAN COLOUR IN THE EASTERN MEDITERRANEAN

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#### Abstract

We study the impact of Saharan dust deposition on chlorophyll concentrations in the Mediterranean using ocean colour products and simulated dust deposition fluxes. Chlorophyll increases follow most dust deposition events but dust fertilization and wind-related effects cannot be disentangled.

Keywords: Surface water fertilization, Dust deposition, Chlorophyll, SeaWiFS, LMDz-INCA.

### Introduction

African dust deposition significantly contributes to soluble phosphorus inputs to Mediterranean surface waters (1), where it is the limiting nutrient, and probably sustains the marine productivity (2). This work aims at investigating this possible fertilization of low productive Mediterranean waters.

### Data and method

SeaWiFS daily level-3 maps (~9 x 9 km²) of chlorophyll concentration ([Chl]) were obtained from NASA (3) for 1998-2000. Only part of the basin is observed every day because of clouds and orbital characteristics.

The atmospheric dust cycle was simulated at a resolution of 1.84° x 2.25°with LMDz-INCA, a 3-D AGCM coupled with chemistry and aerosols (4). We extracted daily integrated dry and wet dust deposition fluxes (F =  $F_{\rm d} + F_{\rm w}$ ).

We averaged [Chl] within the model grid (~600 pixels/cell) and selected four grid cells across eastern Mediterranean: Southern Central (SC: 18°E, 33.06°N), Eastern Central (EC: 22.5°E, 34.90°N), South-western Levantine (SL: 24.75°E, 33.06°N), and Northern Levantine (NL: 29.25°E, 34.90°N). We rejected days with  $\leq\!25\%$  of pixels per grid cell. This left 50-60% of days and we filled gaps using a 7-day moving average.

## Results and discussion

Maximum (resp. minimum) surface [Chl] at a given site varies between 0.20 and 0.34 (0.04 and 0.06) mg m<sup>-3</sup>. The seasonal cycle shows a winter maximum of 0.20±0.05 mg m<sup>-3</sup> in January-February, with another, occasionally absolute, maximum in March (fig.1). [Chl] drops down in spring and remains low (0.05-0.10 mg m<sup>-3</sup>) from May to September. From October, [Chl] slowly increases to its winter maximum.

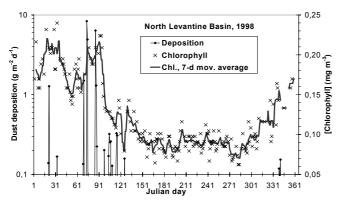


Fig. 1. Daily dust deposition, daily and 7-d averaged chlorophyll at NL in 1998.

A dust outbreak generally causes high deposition at most sites. Overall range in F is 8.9-38.7 g m² yr¹, with a factor 2-3 of interannual variability at a given site, and no clear trend from one year or one site to the other. The fallout is controlled by a few wet deposition events ( $F_w/F > 93\%$ ). Maximum  $F_w$  of 15.7 g m² yr¹ is simulated at SC and  $F_w > 1$  g m² d is reached up to 9 d yr¹.  $F_d$  exceeds 0.1 g m² yr¹ for only 9-10 d yr¹, with a maximum of 1.24 g m² yr¹ at SL. Most of highest deposition events and annual deposition occur during March-May. Summer fallout never exceeds 0.25 g m² d¹. The simulated seasonal cycle is consistent with observations (5, 6).

Laboratory experiments support an increase of 8 mg Chl m<sup>-2</sup> (or 0.08 mg Chl m<sup>-3</sup> in a surface mixed layer of 100 m) per g m<sup>-2</sup> of dust deposition (7) with a lag time of ~48 hours (8). For a summer dust deposition of 0.1 g m<sup>-2</sup> in a mixed layer of 20 m, the expected increase is 0.04 mg Chl m<sup>-3</sup>. Our data set shows in accordance 0.02-0.16 mg m<sup>-3</sup> increases in [Chl] shortly following dust fallout, and thus supports a dust fertilization of Mediterranean waters.

However, we also find a good correspondence between high surface winds and [Chl] peaks, and dust events are generally associated with wind peaks (fig.2). Surface wind controls the mixing between surface waters and deeper ones, richer in nutrients. Observed [Chl] increases can thus be contributed by both dust deposition and wind. Coupled biogeochemical-circulation model and appropriate in situ measurements are necessary to unravel both effects.

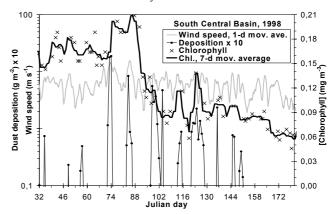


Fig. 2. Daily dust deposition, daily and 7-d averaged chlorophyll, and wind speed at SC, between February 1 and June 30 1998

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# References

- 1 Bergametti G., *et al.*, 1992. Source, transport and deposition of atmospheric phosphorus over the northwestern Mediterranean. *J. Atmos. Chem.*, 14: 501-513.
- 2 Ridame C., and Guieu C., 2002. Saharan input of phosphate to the oligotrophic water of the open western Mediterranean Sea. *Limnol. Oceanog.*, 47: 856-869.
- 3 http://seawifs.gsfc.nasa.gov/SEAWIFS.html
- 4 Bauer S.E., *et al.*, in press. Global modelling of heterogeneous chemistry on mineral aerosol surfaces: The influence on tropospheric ozone chemistry and comparison to observations. *J. Geophys. Res.* 5 Nihlén T., *et al.*, 1995. Monitoring Saharan dust fallout on Crete and its contribution to soil formation. *Tellus*, 47B: 365-374.
- 6 Kubilay N., et al., 2000. An illustration of the transport and deposition of mineral dust onto the eastern Mediterranean. Atmos.
- deposition of mineral dust onto the eastern Mediterranean. *Atmos. Environ.*, 34: 1293-1303.

  7 Ridame C., and Guieu C., 2002. Saharan input of phosphate to
- 7 Ridame C., and Guieu C., 2002. Saharan input of phosphate to the oligotrophic water of the open western Mediterranean Sea. *Limnol. Oceanog.*, 47: 856-869.
- 8 -. Ridame C., 2001. Rôle des apports atmosphériques d'origine continentale dans la biogéochimie marine: Impact des apports sahariens sur la production primaire en Méditerranée. PhD diss., Univ. Paris-6.